

# Formation of atmospheric H<sub>2</sub>SO<sub>4</sub> / H<sub>2</sub>O particles in the absence of organics

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## INTRODUCTION

Atmospheric particles have a strong impact on the Earth's radiation budget due to their radiative properties and the fact that they can act as condensation nuclei for clouds. Field measurements at ground level show atmospheric nucleation events for H<sub>2</sub>SO<sub>4</sub> concentrations of  $\sim 10^7$  molecule cm<sup>-3</sup> (Weber *et al.*, 1999). Despite intensive research activities in the last decade, the mechanism leading to new particles has not been unambiguously revealed yet.

In a previous investigation from this laboratory, experimental evidence for the formation of new particles in the system H<sub>2</sub>SO<sub>4</sub> / H<sub>2</sub>O under near-atmospheric conditions with H<sub>2</sub>SO<sub>4</sub> concentrations of  $\sim 10^7$  molecule cm<sup>-3</sup> was found (Berndt *et al.*, 2005). Here, H<sub>2</sub>SO<sub>4</sub> was produced in situ via the reaction of OH radicals with SO<sub>2</sub> in the presence of water vapor. Similar observations are reported from investigations in the 590 m<sup>3</sup> Calspan chamber (Verheggen, 2004). In both studies organics were present in the reaction gas.

In contrast, taking H<sub>2</sub>SO<sub>4</sub> from a liquid sample  $\sim 10^{10}$  molecule cm<sup>-3</sup> of H<sub>2</sub>SO<sub>4</sub> are needed for producing new particles (Ball *et al.*, 1999).

Subject of this study is to investigate whether the low threshold H<sub>2</sub>SO<sub>4</sub> concentration for nucleation of  $\sim 10^7$  molecule cm<sup>-3</sup> for in situ produced H<sub>2</sub>SO<sub>4</sub> arises from any contributions of organic substances.

## METHODS

The experiments have been performed in the *Ift*-LFT (*Institute for Tropospheric Research* – Laminar Flow Tube; i.d. 8 cm; length 505 cm) at atmospheric pressure and  $293 \pm 0.5$  K using synthetic air as the carrier gas. Particle size distributions ( $dp > 2$  nm) were determined using a differential mobility particle sizer (DMPS) consisting of a short Vienna-type differential mobility analyzer (DMA) and an ultrafine particle counter (UCPC, TSI 3025). From the measured mobility distributions, particle size distributions were determined using an inversion algorithm accounting for the counting efficiency of the UCPC, particle losses inside the DMA and the sampling lines connecting the outlet of the *Ift*-LFT to the DMPS system. In addition, the total number of particles was determined by a second UCPC (TSI 3025) directly at the outlet of the *Ift*-LFT.

H<sub>2</sub>SO<sub>4</sub> was produced in situ via the reaction of OH radicals with SO<sub>2</sub>. In most cases, photolysis of O<sub>3</sub> with the subsequent reaction of formed O(<sup>1</sup>D) with water

served as source of OH radicals. Furan (organic) was added for OH radical titration. The measured furan decay allowed the calculation of the OH radical profile in the *Ift*-LFT and, consequently, the H<sub>2</sub>SO<sub>4</sub> concentration. Simultaneous measurements of the furan and O<sub>3</sub> decay yielded a ratio “reacted O<sub>3</sub>” / “reacted furan” =  $0.5 \pm 0.2$  being in line with the fact that per “effective” photolyzed O<sub>3</sub> molecule two OH radicals were formed. Substituting furan by CO, i.e. removing the organics from the system, the H<sub>2</sub>SO<sub>4</sub> concentrations in the tube were calculated using the measured O<sub>3</sub> decay.

In addition, in few experiments the needed OH radicals were produced via ozonolysis of *t*-butene (dark reaction).

## RESULTS AND DISCUSSION

Figure 1 shows the experimental results for the different approaches for OH radical generation.

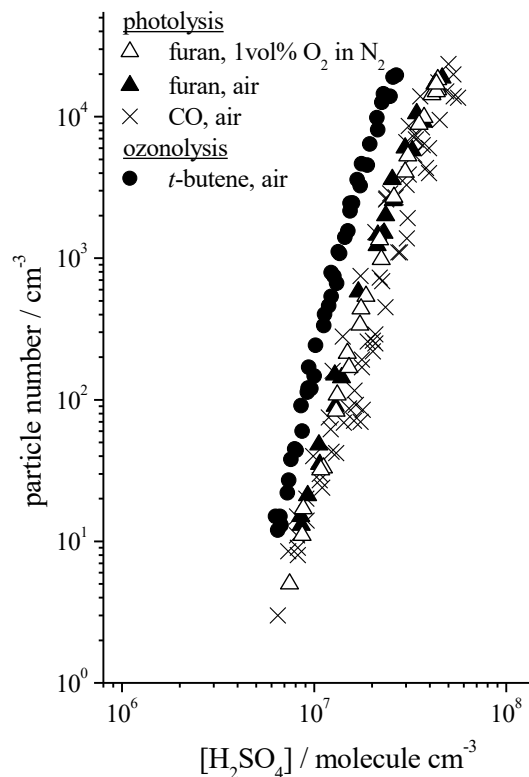


Figure 1. Observed particle numbers (TSI 3025) depending on H<sub>2</sub>SO<sub>4</sub> concentration at 293 K and r.h. = 22 %.

For comparison, in Figure 1 a set of runs ( $O_3$  photolysis in the presence of furan) using 1 vol%  $O_2$  in  $N_2$  as the carrier gas is also added. The findings reveal that the change of the  $O_2$  concentration (1 vol% or 20 vol%) does not influence the particle formation. All four series of measurements consistently demonstrate that new particle formation starts for in situ produced  $H_2SO_4$  concentrations of  $\sim 7 \cdot 10^6$  molecule  $cm^{-3}$ . Therefore, it has to be concluded that organics are not implicitly needed for the process of particle formation under these conditions.

The  $H_2SO_4$  values are affected with an uncertainty of a factor of  $\sim 2$  caused by uncertainties of the measurements needed (reacted furan or  $O_3$ ) or by the simplifications in the chemical mechanisms used for the calculation of  $H_2SO_4$  concentrations. Differences between the four series among each other can be explained in the same way.

Experiments with enhanced  $H_2SO_4$  concentrations allowed the determination of particle size distributions. Figure 2 shows DMPS measurements for three different  $H_2SO_4$  concentrations at r.h. = 42 %.

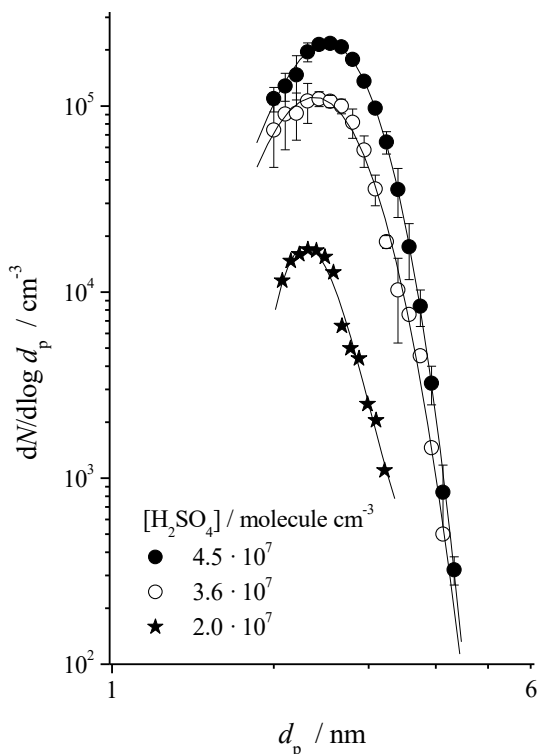


Figure 2. Particle size distributions (DMPS) for three different  $H_2SO_4$  concentrations at r.h. = 42 %. Mean diameters of the particles are 2.5, 2.4, and 2.3 nm for  $H_2SO_4$  concentrations of  $4.5 \cdot 10^7$ ,  $3.6 \cdot 10^7$ , and  $2.0 \cdot 10^7$  molecule  $cm^{-3}$ , respectively. OH radical formation via  $O_3$  photolysis in the presence of CO.

Even for the highest  $H_2SO_4$  concentrations, the mean particle diameter was below 3 nm. A particle diameter of  $\sim 2$  nm for an  $H_2SO_4$  concentration of  $10^7$  molecule  $cm^{-3}$  can be roughly estimated from the observed trend of decreasing particle diameter with decreasing  $H_2SO_4$  concentration. Assuming that the particles consist of  $H_2SO_4$  only (density:  $1.85$  g  $cm^{-3}$ ), the upper limit of the yields of particulate  $H_2SO_4$  formed from gaseous  $H_2SO_4$  can be given. According to this, a yield of 0.9, 5, and 9 % was found for  $H_2SO_4$  vapor concentrations of  $2.0 \cdot 10^7$ ,  $3.6 \cdot 10^7$ , and  $4.5 \cdot 10^7$  molecule  $cm^{-3}$ , respectively, cf. Figure 2.

The results from our DMPS measurements were confirmed by simultaneous runs using a Nano-SMPS system (TSI 3936-N86 with TSI 3786).

## CONCLUSIONS

The results from this study show that the low threshold concentration for nucleation of  $\sim 10^7$  molecule  $cm^{-3}$  in the case of in situ produced  $H_2SO_4$  (from bulk sample  $\sim 10^{10}$  molecule  $cm^{-3}$  of  $H_2SO_4$ ) cannot be explained by a contribution of organics supporting the particle formation. Probably, in the course of the  $SO_2$  oxidation ( $SO_2 \rightarrow SO_3 \rightarrow H_2SO_4$ ) other species than the expected  $H_2SO_4$ (-hydrates) are formed being responsible for particle formation. That is highly speculative at present and much more work is needed for clarifying the nucleation mechanism.

**Keywords:** Atmospheric Aerosol, Nucleation,  $H_2SO_4$

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