

Influence of NH₃ on atmospheric particle formation starting from OH + SO₂

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Introduction

For more than a decade the formation of new aerosol particles in the atmosphere has been the subject of intense studies in the field and in the laboratory. The formation mechanisms and the participating substances, however, have not been resolved yet (Kulmala, 2003). Large discrepancies between model-predicted nucleation rates for the favored binary system H₂SO₄ / H₂O and much higher atmospheric nucleation data were explained by various supportive additional participants such as NH₃ or organic molecules.

In a previous investigation of our laboratory under near-atmospheric conditions, experimental evidence for the formation of new particles in the system H₂SO₄ / H₂O was found for H₂SO₄ concentrations of 10⁷ molecule cm⁻³ if "H₂SO₄" was produced in-situ via the reaction of OH radicals with SO₂ ("H₂SO₄" stands for all products of converted SO₂), Berndt et al. (2005). In contrast, taking H₂SO₄ from a liquid reservoir concentrations of 10⁹ - 10¹⁰ molecule cm⁻³ are needed for new particle formation (Ball et al., 1999, Zhang et al., 2004). The accepted mechanism of atmospheric SO₂ oxidation is as follows:



In competition to pathway (2) HOSO₂O₂ can be formed via:



It can be hypothesized that HOSO₂O₂ or its hydrated species HOSO₂O₂(H₂O)_n play a role for the nucleation process explaining the different threshold concentrations for nucleation starting from liquid H₂SO₄ or from OH + SO₂.

The activities focused on experimental studies regarding the influence of NH₃ on new particle formation. Theoretical studies predict strong enhancement of the nucleation rate for NH₃ mixing ratios at pptv-level (Coffman and Hegg, 1995). Ball et al. (1999) showed experimentally an increase of the nucleation rate for NH₃ mixing ratios of 0.08 – 0.17 ppbv and a relative humidity in the system of 5 - 15 %. In their study H₂SO₄ from a liquid reservoir was used. Subject of this experimental study is the role of NH₃ in the process of new particle formation starting from OH + SO₂. A comparison with atmospheric measurements includes also recent results obtained in the absence of NH₃ additions.

Results in absence of NH₃

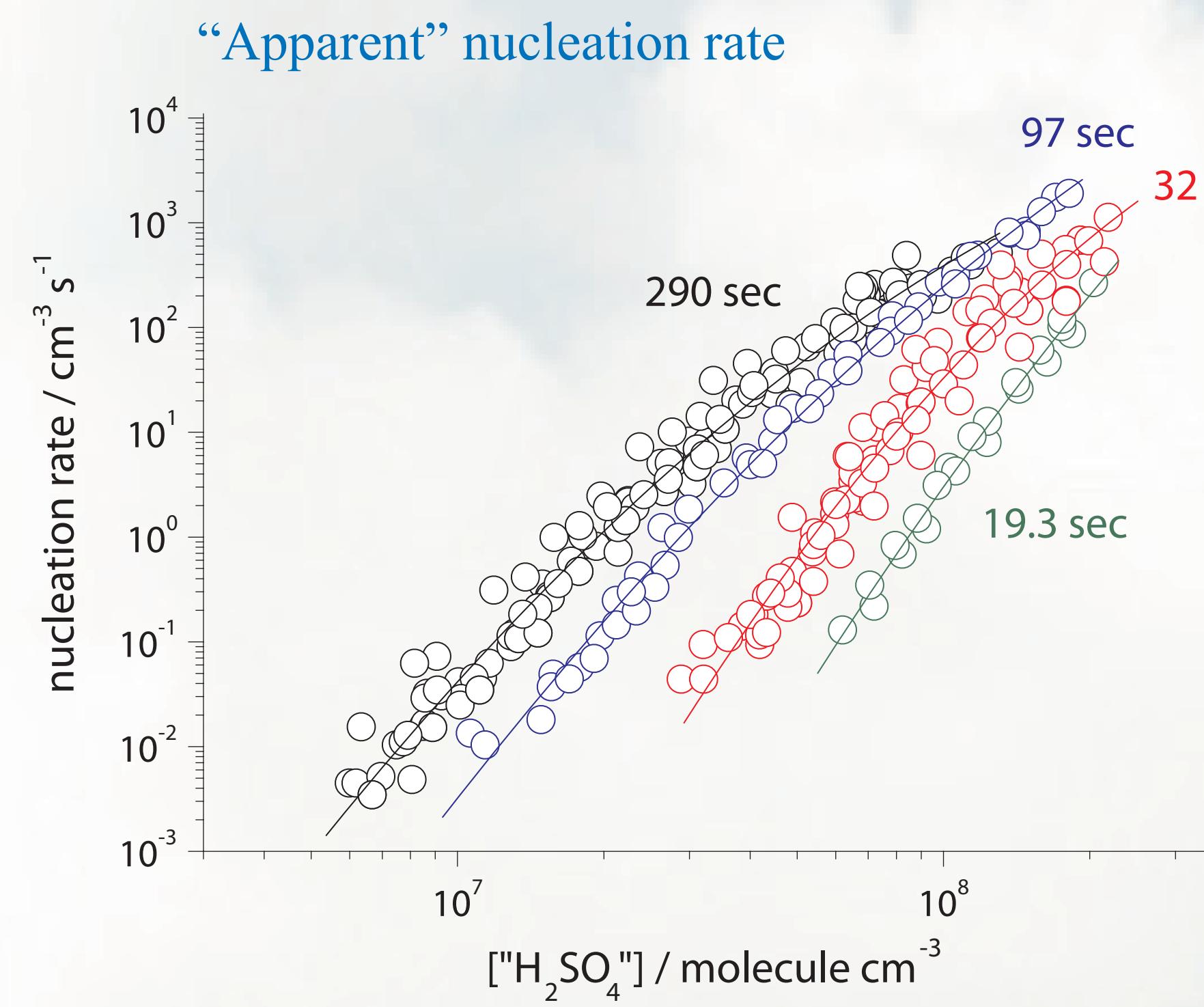


Figure 1: "Apparent" nucleation rate as a function of "H₂SO₄" concentration for different residence times in the flow tube. The nucleation rate is obtained by dividing the measured total particle number by the residence time in the irradiated middle section, r.h.: 22%, TSI 3025.

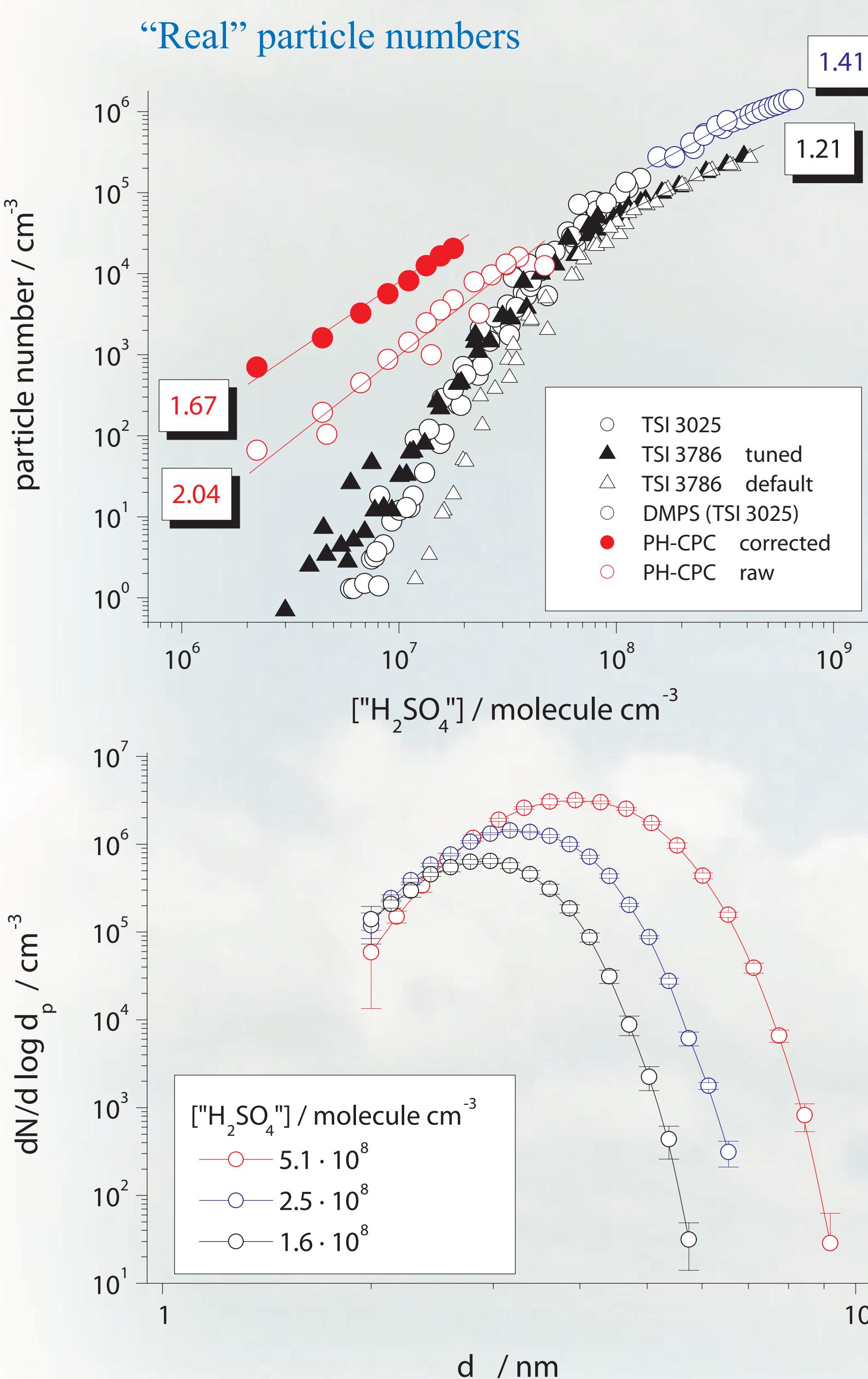


Figure 2: Measured total particle number using three different types of CPCs for integral measurements as well as from DMPS. Given numbers represent the slopes N vs. "H₂SO₄", (TSI 3786: N > 50.000 cm⁻³; DMPS: N < 750.000 cm⁻³), r.h.: 22%, total flow rate: 3.33 standard litre min⁻¹ (residence time in irradiated middle section: 290 sec).

Figure 3: Examples of DMPS size distributions recorded for experimental conditions as given in figure 2. In the case of PH-CPC measurements, simulations yield mean particle diameters in the range of 1.5 - 2 nm.

Experiments have been performed in the atmospheric pressure flow-tube Ift-LFT (i.d. 8 cm; length 505 cm) at 293 +/- 0.5 K. The first tube section (56 cm) includes an inlet system for gas entrance (humidified air premixed with CO, SO₂, and O₃). The middle section (344 cm) is equipped with 8 UV lamps for a homogeneous irradiation. A non-irradiated end section (105 cm) holds the sampling outlet. Gas-phase analysis was carried out using analyzers for O₃ and SO₂ (Thermo Environmental Instruments: 49C and 43C), a butanol-based UCPC (TSI 3025), a H₂O-based UCPC (TSI 3786) as well as pulse height CPC for integral particle measurements, and a differential mobility particle sizer (Vienna-type DMA with UCPC, TSI 3025) for monitoring of size distributions. The pulse height CPC (PH-CPC) comprises a TSI-3025A ultrafine CPC with modified optics and a multichannel analyzer, e.g. Weber et al. (1998), allowing particle detection for sizes well below 2 nm in mobility diameter (Sipilä et al., 2008). NH₃ concentrations in the carrier gas were measured before entering the flow tube as well as at the outlet by means of a trace gas monitor TGA 310 (OMNISENS) with a stated detection limit of 2.5 · 10⁹ molecule cm⁻³.

Experimental setup



Head of the flow tube: Non-irradiated first section with the gas inlets as well as the gas mixing devices

Determination of ["H₂SO₄"]



The effective photolysis rate coefficient k₄ was determined in each experiment measuring the O₃ decay. With the knowledge of k₄ along with k₅ – k₇ taken from literature as well as the concentrations for O₃, CO, and SO₂ the resulting "H₂SO₄" concentration in the tube was calculated solving the resulting differential equations. All experiments have been done in the absence of organics.

Results in presence of NH₃

"Real" particle numbers

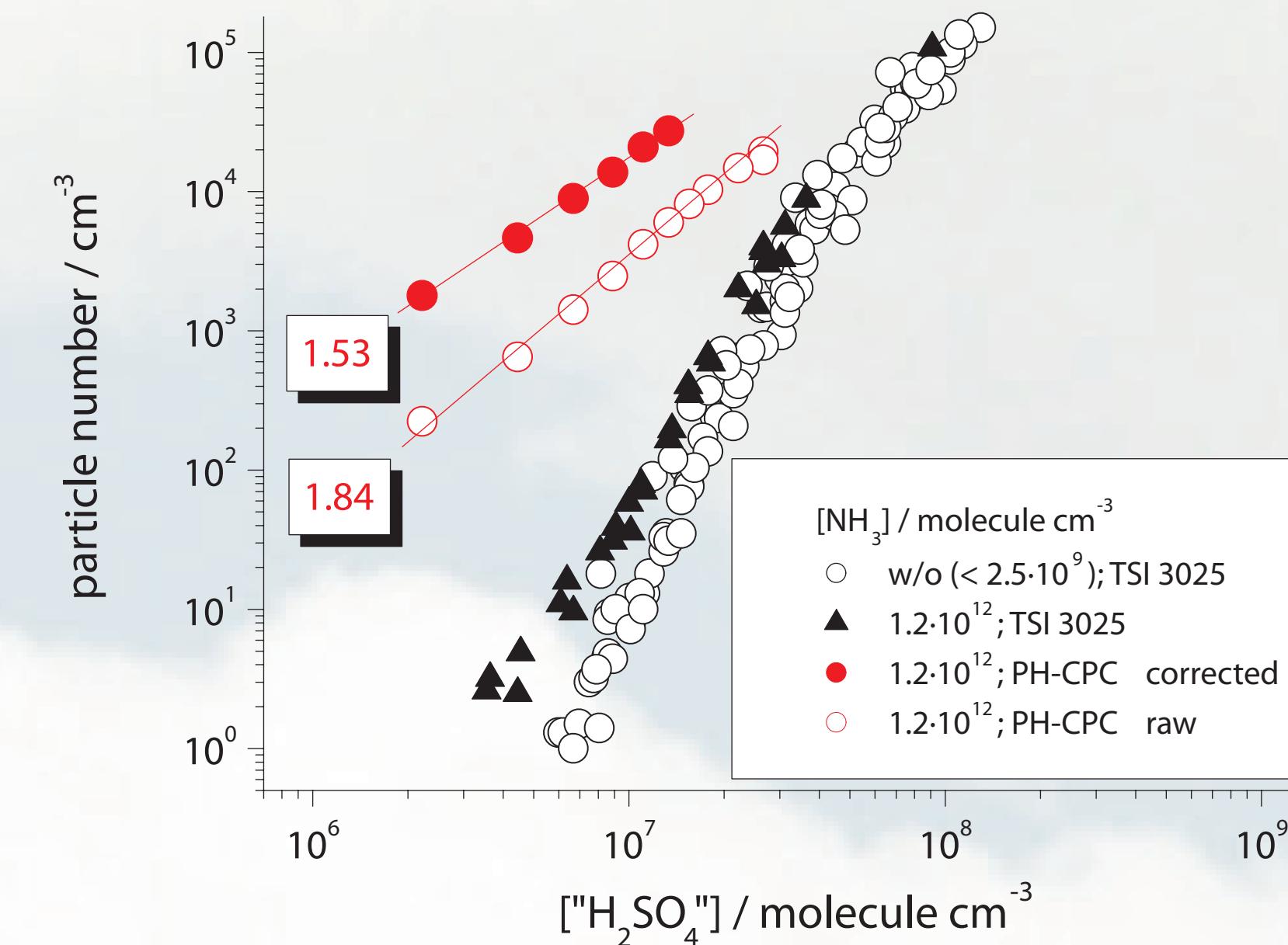


Figure 4a: Total particle numbers vs. "H₂SO₄" concentration in absence and presence of NH₃ using TSI 3025 and PH-CPC, r.h.: 22%, total flow rate: 3.33 standard litre min⁻¹ (residence time in irradiated middle section: 290 sec).

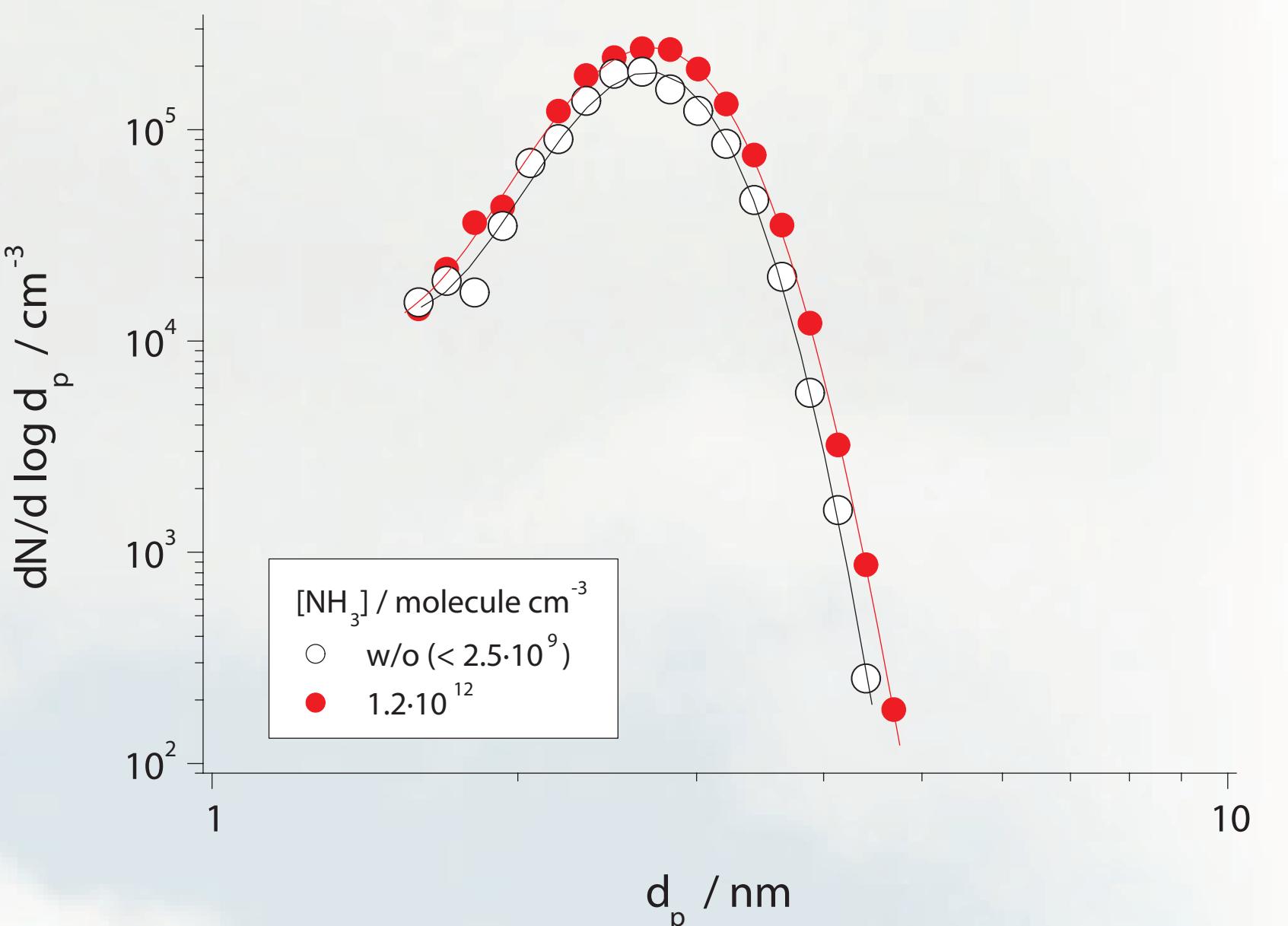


Figure 4b: Size distributions measured in absence and presence of NH₃ for a "H₂SO₄" concentration of 9 · 10⁷ molecule cm⁻³, cf. conditions as given in figure 4a.

Influence of relative humidity

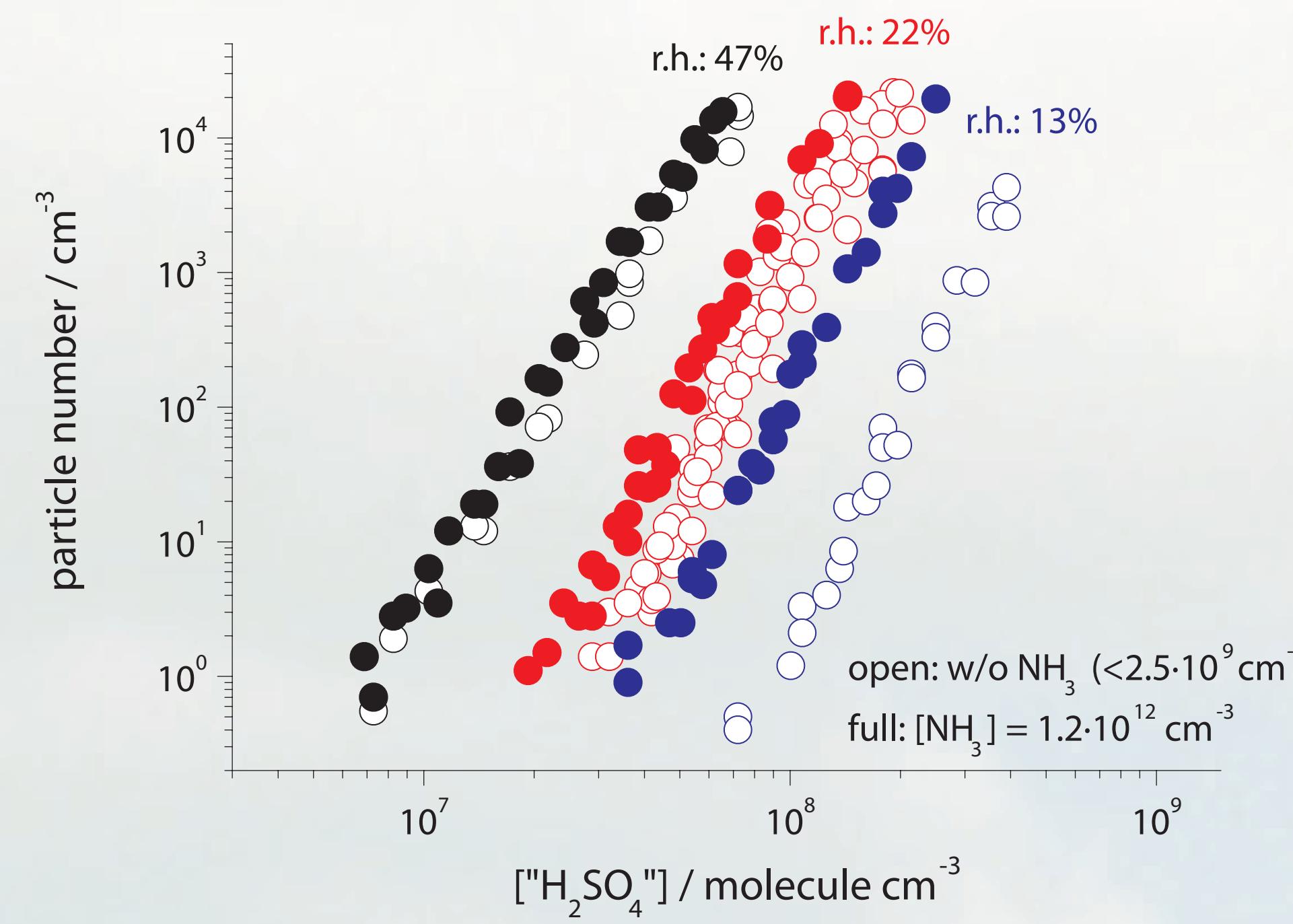
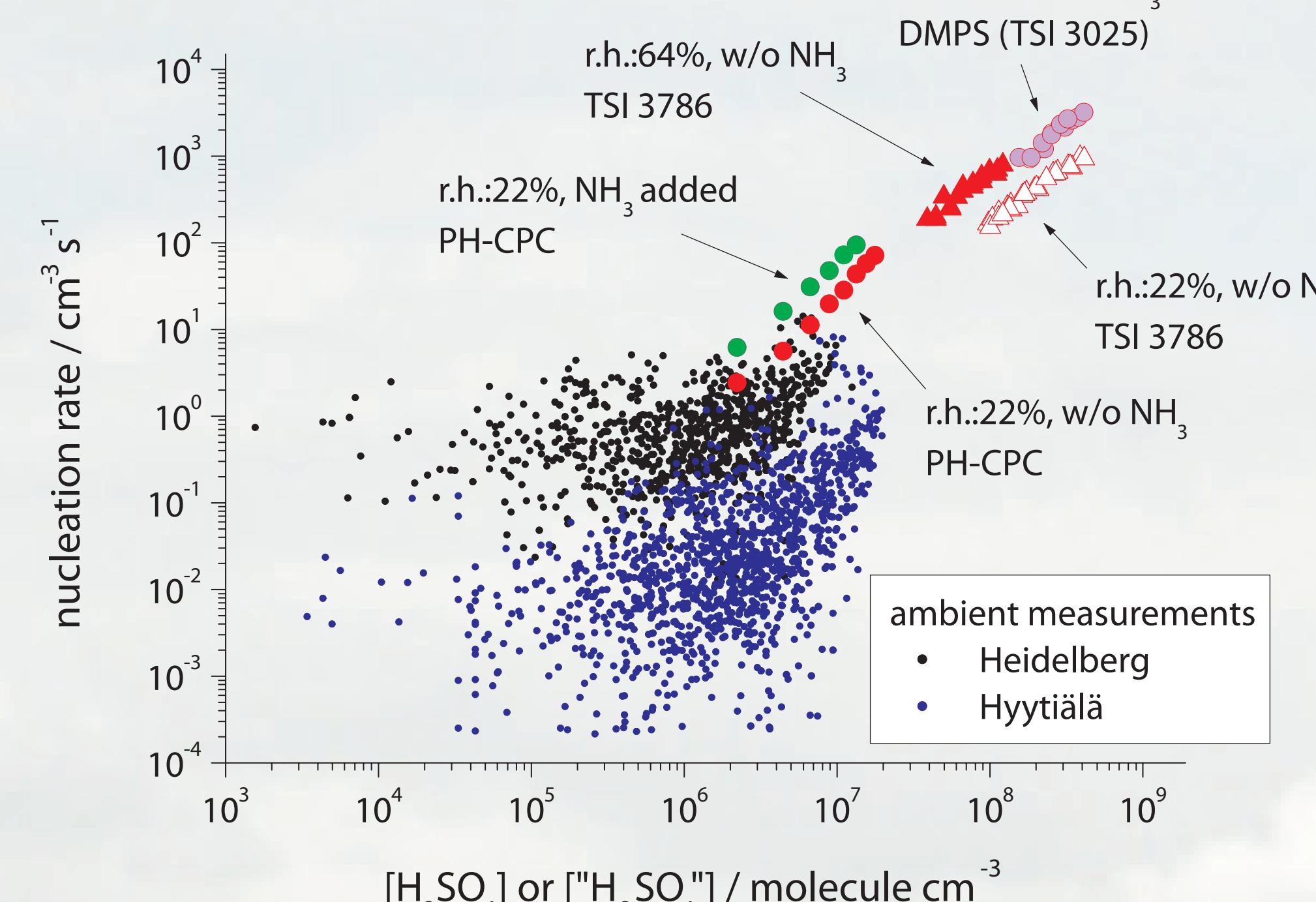


Figure 5: Total particle numbers vs. "H₂SO₄" concentration in absence of NH₃ (< 2.5 · 10⁻⁹ molecule cm⁻³; outlet: 1.1 · 10⁻¹² molecule cm⁻³) for three different relative humidities, total flow rate: 30 standard litre min⁻¹ (residence time in irradiated middle section: 32 sec) TSI 3025.

In agreement with former results, a distinct increase of total particle number with increasing r.h. was observed for the measurements performed in the absence of NH₃. NH₃ addition increased the total particle number by about one order of magnitude at r.h. = 13%, by a factor of 3 – 5 at r.h. = 22% and by a factor of less than two at r.h. = 47%. From mechanistic point of view it is highly speculative to explain this r.h. dependence. A possible reason could be the competing steps i) reaction of SO₃ with 2 water molecules (or with the dimer) producing H₂SO₄ and ii) reaction of SO₃ with NH₃ producing sulfamic acid. Sulfamic acid can significantly contribute to particle formation and growth.

Application to the atmosphere



Laboratory:
slope log N vs. log ["H₂SO₄"] = 1 - 2
(nearly independent on NH₃ addition)

Figure 6: Comparison of J vs. [H₂SO₄] or ["H₂SO₄"] data measured either in lower atmosphere or in laboratory. Ambient measurements from Heidelberg (28.2. - 3.4.2004) and from Hyttiala (5.4. - 16.5.2005) were adopted from Riipinen et al. (2007). Ambient data can be described with a power-law dependence on H₂SO₄ with typical powers of 1 - 2, cf. Riipinen et al. (2007).

References:

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