

Measurements of secondary organic aerosol formation by aerosol mass spectrometry during ozonolysis of terpenes

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Introduction

Aerosol plays a key role in atmospheric processes. They can directly or indirectly influence the radiative forcing, decrease the visibility, and affect the human health. Secondary organic aerosol (SOA) is formed from atmospheric reactions of both anthropogenic and biogenic volatile organic compounds. Among these biogenic SOA precursors, it is known that gas phase reactions of monoterpenes represent an important source of biogenic SOA. However, knowledge on mechanisms of formation and the variation of the chemical composition of monoterpenes SOA is still relatively sparse. Thus, the formation of SOA during the ozonolysis of three different monoterpenes (limonene, α -pinene and β -pinene) was investigated in an aerosol chamber. Formation of SOA was followed by measuring vacuum aerodynamic size distribution and chemical composition using an Aerodyne Aerosol Mass Spectrometer.

Aerosol Chamber

A series of monoterpane ozonolysis experiments was carried out in the LEAK chamber (Leipzig Aerosol Kammer). This chamber has a volume of 17 m^3 and a surface to volume ratio (S/V) of 2.1. Temperature, ozone concentration and relative humidity were monitored during the course of experiment. Seed particles were generated by atomizing an acidic solution of ammonium sulfate ($0.03\text{M}/0.05\text{M}$ $(\text{NH}_4)_2\text{SO}_4/\text{H}_2\text{SO}_4$). Parallel to the AMS measurements, the particle size distribution was monitored by a SMPS system.

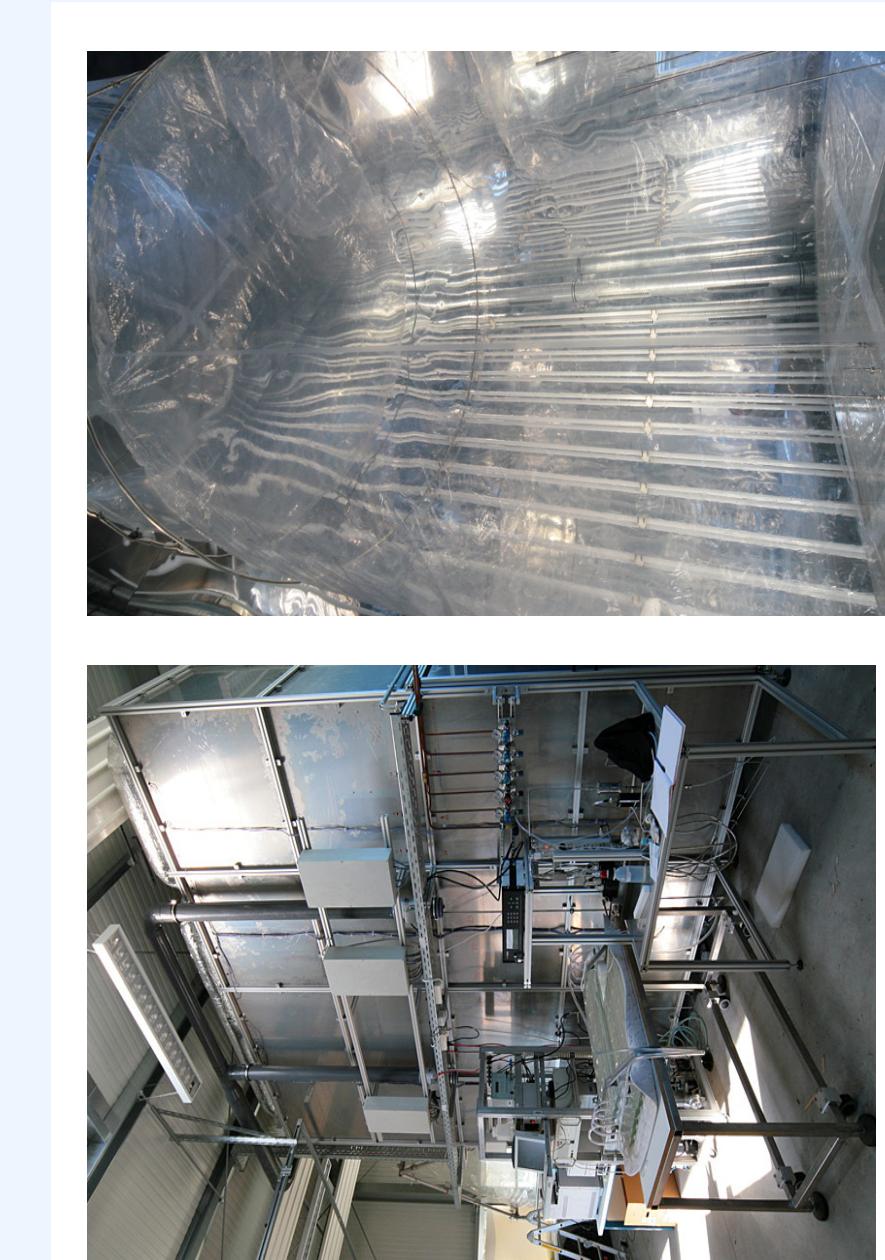


Figure 1: LEAK aerosol chamber at the ITT

Results

SOA formation

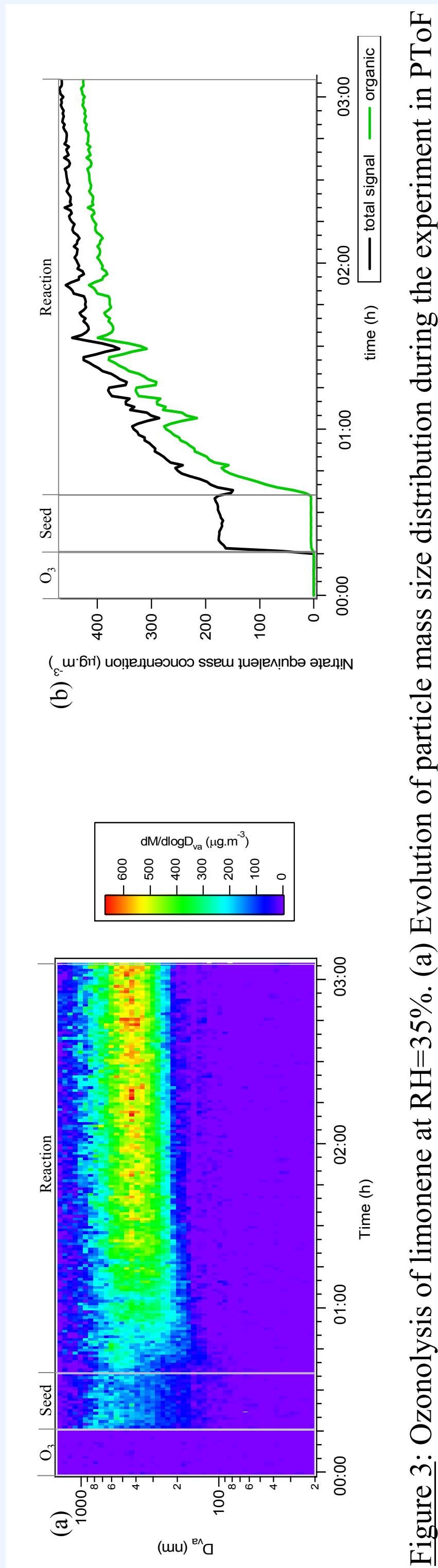


Figure 3: Ozonolysis of limonene at RH=35%. (a) Evolution of particle mass size distribution during the experiment in PToF mode and (b) time profile concentration of total signal and organics in MS mode.

Impact of relative humidity

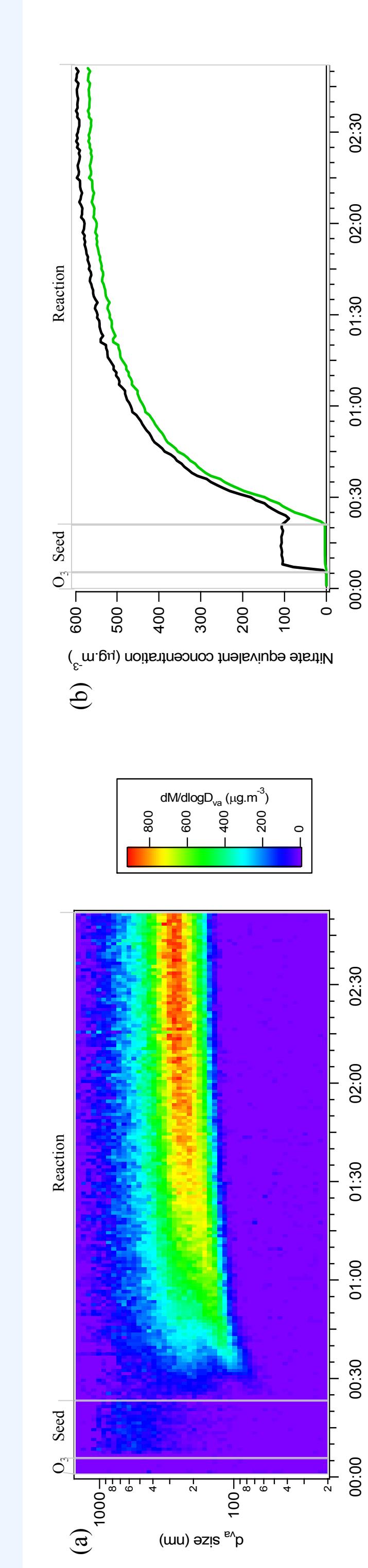


Figure 4: Ozonolysis of limonene at RH<10%. (a) Evolution of particle mass size distribution during the experiment in PToF mode and (b) time profile concentration of total signal and organics in MS mode.

AMS

Formation of SOA was followed by an Aerosol High Resolution Time of Flight Mass Spectrometer (HR-ToF-AMS) that provides real-time size resolved aerosol chemical composition (Jayne et al. 2000; Canagaratna et al. 2007). Particles are focused into a narrow beam for “Time of Flight” mode measurements (i.e. P-TOF mass distribution measurement mode) or to block/clear the particle beam for “mass spectrum” mode measurement (MS mode). The non-refractory components of the particles are flash vaporized on a heated surface (600°C) at the end of the P-TOF chamber. The evaporated particles are ionized by electron impact source (70eV). Positive ions are then extracted into the time of flight mass spectrometer with a frequency of 80Hz for their mass-to-charge-ratio (m/z) determination. During this work, the mass spectrometer was only used in V-mode.



Figure 2: Aerodyne Aerosol High Resolution Time of Flight Mass Spectrometer (HR-ToF-AMS)

Experimental conditions

Initial conditions are summarized in Table 1. In all cases the ratio between the reactant and ozone was chosen in relation with the rate constant of each monoterpe. Before each experiment, chamber is flushed by clean air at the corresponding relative humidity for 16h. Subsequently chamber is closed and ozone is introduced. Seed particles are added into the chamber when ozone concentration is reached at the set level.

The concentration of seed particles was measured for 10-15 min. This allowed us to obtain an accurate initial seed particle concentration. At last, the reactant is introduced into the chamber by vaporizing a pure solution. In all experiments, AMS measurements started during the ozone injection. This period is totally free from organic compounds and these measurements are considered as blank for the AMS signals.

References

- Alfarra M.R., Paulsen D., Gysel M., Garforth A.A., Dommen J., Prevôt A.S.H., Worsnop D.R., Baltensperger U. and Jimenez J.I., Atmospheric Chemistry and Physics, 6, 5279-5293 (2006).
- Bahreini R., Keywood M.D., Ng N.L., Vaattabangul V., Gao S., Flagan R.C., Jimenez J.I., Worsnop D.R. and Jimenez J.I., Environmental Science and Technology, 39, 5674 (2005).
- Canagaratna M.R., et al., Mass Spectroscopy Reviews, 26, 185 (2007).
- DeCarlo P.F., Slowik J.G., Worsnop D.R., Jimenez J.I., Aerosol Science and Technology, 38 (12), 1185 (2004).
- Jayne J.T., Lear D.C., Davidovits P., Smith K.A., Kolb C.E. and Worsnop D.R., Aerosol Science and Technology, 33, 49 (2000).
- Takegawa N., Miyakawa T., Kawamura K., Kondo Y., Aerosol Science and Technology, in press.

Summary

The formation of SOA during ozonolysis of limonene, α -pinene and β -pinene was followed by AMS in presence of acidic ammonium sulfate seed particles. In all cases, the organic fraction of the particle rapidly increased after the beginning of the reaction. Impact of the relative humidity was investigated for limonene. The formation of smaller particles in the first steps of the reaction and the presence of more organic compounds at the end of experiment than at RH=35% were observed. The initial concentration of seed particles was slightly lower in the dry condition that might have also contributed to this differences. The MS mode showed that the carbonyl (m/z 43) and oxo- and dicarboxylic acids (m/z 44) functions represented the most important component of organic fragments in all cases. These indicate that the SOAs are composed of highly oxidized compounds. The presence of higher molecular weight compounds was also observed.

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